ATSC 201 Fall 2024 Chapter 4: A1f, A4f, A13f, A17f, A25f, E17

Total marks out of 40

## **Chapter 4**

A1f) (5 marks)

Compare the saturation vapor pressures (with respect to liquid water) s) calculated with the Clausius-Clapeyron equation and with Tetens' formula, for T(°C): e) 20

Given: T = 20 °C 293.15 K

Use eqn 4.1a (Clausius Clapeyron eqn)

 $e_{s} = e_{o} \cdot \exp\left[\frac{L}{\Re_{v}} \cdot \left(\frac{1}{T_{o}} - \frac{1}{T}\right)\right]$ 

where Rv =	461	J/(K*kg)
To =	273.15	К
eo =	0.611	kPa
L = Lv (for liqui	d water) =	2.50E+06 J/kg

Use eqn 4.2 (Teten' formula)

 $e_s = e_o \cdot \exp\left[\frac{b \cdot (T - T_1)}{T - T_2}\right]$ 

where b =	17.2694
eo =	0.611 kPa
T1 =	273.15 K
T2 =	35.86 K

Clausius-Clapeyron:

es = 2.368 kPa

Teten's:

es = 2.339 kPa

Clausius-Clapeyron's saturation vapor pressure is slightly larger than Teten's.

Check: Units ok. Physics ok.

Discussion: The Clausius-Clapeyron gives a saturation vapor pressure slightly

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larger to Teten's formula for a temperature of 25°C. This is consistent with the graph on p.89 from the sample application. Teten's formula accounts for L varying slightly, but it appears the difference is usually negligible.

A4f) (18 marks)	Calculate the values of es (kPa), r (g/kg), q (g/kg), ρv (g/m^3), RH (%), Td (°C), LCL (km), Tw (°C), rs(g/kg), qs (g/kg), and ρvs (g/m^3), given the following atmospheric state: ef) P = 90kPa, T = 5°C, e = 0.2kPa.						
	Given:	P = T = e =		90 kPa 5 degC 0.2 kPa	278.15 K		
	Find:	es = r = φv = RH = Td = LCL = Tw = rs = φvs =	? ? ? ? ? ? ? ? ? ? ? ? ? ? ? ? ? ? ?	kPa g/kg g/m^3 % degC km degC g/kg g/kg g/kg			
	Use eqn. 4.1	La:	$e_s = e_0 \cdot \exp\left[$	$\left[\frac{L}{\Re_v} \cdot \left(\frac{1}{T_o} - \frac{1}{T}\right)\right]$			
	where eo = To =		0.611 kPa 273.15 K				
	Lv/Rv =		5423 K				
	es =		0.87 kPa				
		• r - e*	(P - 0)				

Use eqn 4.4:  $r = \epsilon^* e / (P - e)$ where  $\epsilon = 622 \text{ g/kg}$  0.622 g/g

Use eqn. 4.7:  $q = \epsilon^* e/P$ 

q = 1.38 g/kg

Use eqn. 4.10:  $\rho v = e/(Rv * T)$ where  $Rv = 4.61E-04 \text{ kPa*m^3/(K*g)}$ 

Use eqn. 4.14a: RH =( e/ es) \* 100

RH =	22.91 %
Use eqn. 4.15a:	$T_d = \left[\frac{1}{T_o} - \frac{\Re_v}{L} \cdot \ln\left(\frac{e}{e_o}\right)\right]^{-1}$
where eo =	0.611 kPa
To =	273 K
Rv/Lv =	1.84E-04 /K
Td =	258.47 K
Td =	-14.68 degC

Use eqn. 4.16a: LCL = a\*(T-Td) where a = 0.125 km/degC

LCL = 2.46 km
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To find wet bulb temperature using Normand's Rule: Use eqn 4.21: Tw = T\_LcL + Γs\*LCL

Find I s with eqn 4.37a:

$$\Gamma_{s} = \frac{|g|}{C_{p}} \cdot \frac{\left(1 + \frac{r_{s} \cdot L_{v}}{\Re_{d} \cdot T}\right)}{\left(1 + \frac{L_{v}^{2} \cdot r_{s} \cdot \varepsilon}{C_{p} \cdot \Re_{d} \cdot T^{2}}\right)}$$

Use eqn. 3.3: cp = cpd\*(1+1.84\*r) where cpd = 1004 J/(kg\*K)

Tw =	-5.26 degC	
Tw =	267.89 K	_
T_LcL =	254.04 K	
Find T_LcL w where Fd =	ith eq 4.20: T_LcL = T - Гd 9.8 K/km	*LCL
Гs =	5.63 K/km	
Lv/Rd =	8707.77 K	to input)
Lv/cp =	2483.71 K	values to make the huge egn nicer
g/cp =	9.74 K/km	(I calculated some intermediate
cp =	1006.55915 J/(kg*K)	
r (g/g) =	0.00139	

Use eqn 4.5: rs =  $(\epsilon^* es)/(P-es)$ 

rs =	6.09 g/kg
rs =	0.00609279 g/g

Use eqn. 4.8: qs =  $(\epsilon^* es)/P$ 

Use eqn. 4.12: ρvs = es/(Rv\*T)

ρνs = 6.81 g/m^3

Check: Units ok. Physics ok.

Discussion:

Because the air at this level is unsaturated, any liquid

droplets in the air will tend to evaporate. Imagine leaving a wet towel in a room; it will eventually dry up due to evaporation. This evaporation causes cooling in the wet towel. The wet-bulb temperature represents the temperature of the wet towel after the evaporative cooling. Therefore, you would always expect the wet-bulb temperature to be cooler than the dry-bulb temperature in unsaturated air.

<b>A13f)</b> (3 marks)	<ul> <li>For air at sea level, find the total water mixing ratio for a situation where:</li> <li>f) r= 4 g/kg, rL = 1 g/kg.</li> </ul>			where:		
	Given:	r = rL =		4 g/kg 1 g/kg		
	Find:	rT =	?	g/kg		
	Use eqn 4	.32a: rT = rs +	rL + ri			
	Need ri: Assume ri	= 0 g/kg				
	rT =		5 g/kg			

Check: Units ok. Physics ok.

**Discussion:** Since we assume there is no ice presents in our parcel, the amount of water vapor in the air is simply the sum of the water vapor and the liquid present in our parcel.

A17f) (3 marks) Given an air parcel starting at 100kPa with dewpoint (degC) given below, use Fig. 4.7 to find the parcel's final dewpoint (degC) if it rises to a height where P = 60kPa. f) 15



See attached Fig. 4.7.

Tdf = 8 degC

Check: Units ok. Physics ok.

Discussion: The dew point temperature decreases with height because air expands and cools adiabatically - just like with regular temperature.

A25f) (3 marks) Given an air parcel that starts at a height where P = 100kPa with T = 25degC and r = 12 g/kg (ie. it is initially unsaturated). After rising to its final height it has an rL (g/kg) value listed below. Assuming no precipitation falls out, find the value for r (g/kg) for this now-saturated air parcel. f) 3

Given:	P_initial =		100 kPa
	T_initial =		25 degC
	r_initial =		12 g/kg
	rL_final =		3 g/kg
	rL_initial =		0 g/kg
Find:	r_final =	?	g/kg

Use eqn 4.35b:

 $(r + r_I)_{initial} = (r + r_I)_{final} \qquad \bullet (4.35b)$ 

Re-arrange: r\_final = r\_initial + rL\_initial - rL\_final

Check: Units ok. Physics ok.

Discussion: In this situation, we are assuming that the water vapor condenses

into liquid form, and the liquid droplets are suspended in the air. This way, we can assume that the total mixing ratio remains constant, as water is not lost through precipitation.

Create a thermo diagram using a spreadsheet to calculate isohumes (for r = 1,

## E17) (8 marks)

3, 7, 10, 3 the same	30 g/kg) graph v	and dry adiaba 's P on an invert	ts (for $\theta = -30$ , ted log-scale sin	-10, 10, 30 °C), a milar to Figs. 3.3	Ill plotted on 3 and 4.7.	
Find:	Р	lot isohumes	on a thermo	o diagram.		
		[1 94	( r.P	)] <sup>-1</sup>		
Use:		$T_d = \left[\frac{1}{T_o} - \frac{3}{T_o}\right]$	$\frac{v_0}{L} \ln \left( \frac{1}{e_0 \cdot (r+1)} \right)$	<u>ε)</u>	eqn 4.15b	
Or:		$T = \left[\frac{1}{T_o} - \frac{\Re_z}{L_v}\right]$	$\frac{1}{2} \cdot \ln \left\{ \frac{r_s \cdot P}{e_0 \cdot (r_s + r_s)} \right\}$	$\left[\frac{1}{\epsilon}\right]^{-1}$	eqn 4.36	
eo =		0.611 k	¢Pa			
Rv/Lv =		1.84E-04 1	ι/κ			
To =		273 k	<			
e =		0.622 g	3/g			
P (kPa)	Т	for r=1 1	۲ for r=3	T for r=7	T for r=10	T for r=30
	10	-42.27	-30.99	-21.56	-17.39	-3.84
	20	-35.26	-23.27	-13.21	-8.75	5.75
	30	-30.96	-18.52	-8.06	-3.43	11.68
	40	-27.81	-15.03	-4.28	0.49	16.05
	50	-25.31	-12.27	-1.28	3.60	19.52
	60	-23.23	-9.96	1.23	6.19	22.43
	70	-21.44	-7.98	3.38	8.43	24.93
	80	-19.88	-6.24	5.27	10.39	27.13
	90	-18.48	-4.68	6.97	12.15	29.10
	100	-17.21	-3.28	8.50	13.73	30.89



Check: Units ok. Physics ok. Looks like fig 4.7.

**Discussion:** When we are given the air temperature and the dew point temperature of a parcel, we can use the thermodynamic diagram to find its lifting condensation level by following the dry adiabat from the temperature point, and the isohume from the dew-point temperature point.